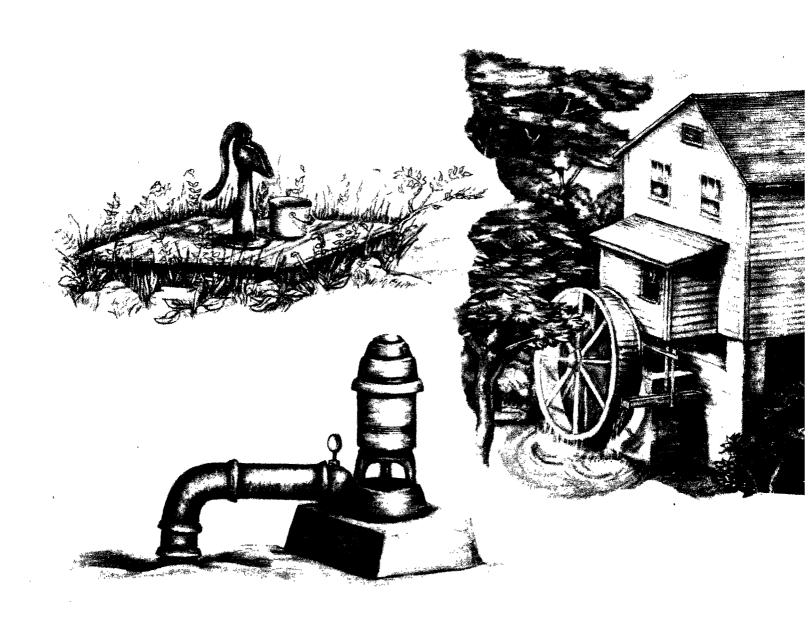


PRELIMINARY

DELINEATION AND DESCRIPTION OF THE REGIONAL AQUIFERS OF TENNESSEE-THE EAST TENNESSEE AQUIFER SYSTEM



Prepared by
U.S. GEOLOGICAL SURVEY
in cooperation with the
U.S. ENVIRONMENTAL PROTECTION
AGENCY

PRELIMINARY DELINEATION AND DESCRIPTION OF THE REGIONAL AQUIFERS OF TENNESSEE--THE EAST TENNESSEE AQUIFER SYSTEM

By John V. Brahana, Dolores Mulderink, Jo Ann Macy, and Michael W. Bradley

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UNITED STATES DEPARTMENT OF THE INTERIOR

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CONVERSION FACTORS

In this report, figures for measures are given only in inch-pound units. Factors for converting inch-pound units to International System of units (SI) are shown in the following table:

| Multiply | <u>B</u> <u>y</u> | To obtain |
|---|-------------------|--|
| inch (in.) | 25.4 | millimeter (mm) |
| foot (ft) | 0.203 | meter (m) |
| cubic foot (ft ³) | 0.02832 | cubic meter (m ³) |
| mile (mi) | 1.609 | kilometer (km) |
| gallon (gal) | 3.785 | liter (L) |
| gallon per minute (gal/min) | 0.0631 | liter per second (L/s) |
| foot per day (ft/d) | 0.305 | meter per day (m/d) |
| feet squared per day (ft ² /d) | 0.0929 | meters squared per day (m ² /d) |

National Geodetic Vertical Datum of 1929 (NGVD of 1929): A geodetic datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called mean sea level. NGVD of 1929 is referred to as sea level in this report.

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ABSTRACT

The East Tennessee aquifer system occurs in the Valley and Ridge and the Blue Ridge provinces of Tennessee. These areas are underlain by rocks of Precambrian to Mississippian age which have been structurally deformed and faulted during the Appalachian orogeny. Ground water in the Valley and Ridge occurs primarily in solution openings in carbonate rocks and in fractures in sandstones and shale. Fractures in the crystalline rocks store and transmit most of the ground water in the Blue Ridge province.

The East Tennessee aquifer system is important as a source of rural and municipal drinking water. Within 300 feet of land surface, ground water generally contains less than 500 milligrams per liter dissolved solids. At greater depths, fractures and solution openings are smaller and fewer in number. There are very few data to define ground-water occurrence at depths greater than about 300 feet. Ground-water flow may be restricted and the dissolved-solids concentrations in the ground water may reach thousands or even tens of thousands of milligrams per liter.

INTRODUCTION

The Safe Drinking Water Act (Public Law 93-523) includes provisions for the protection of underground sources of drinking water. Specifically, Part C of the Act authorizes the Environmental Protection Agency to establish regulations to insure that underground injection of contaminants will not endanger existing or potential sources of drinking water. As developed by EPA, the regulations require that all underground sources of ground water with less than 10,000 milligrams per liter (mg/L) dissolved solids which do not contain hydrocarbon, mineral, or geothermal resources be designated for protection whether they are or are not currently being used as a source of drinking water.

The geologic formations of Tennessee (Miller, 1974) have been delineated on a regional basis into eight major regional aquifers having broad areal extent. Each regional aquifer is characterized by a unique set of hydrologic conditions and water quality.

The purpose of this report is to describe the formations that comprise the East Tennessee aquifer system (fig. 1) and to delineate zones within this aquifer system that are actual or potential drinking-water sources.

This report on the East Tennessee aquifer system provides generalized information on (1) the areal and stratigraphic occurrence of the aquifer, (2) dissolved-solids content of the ground water, (3) area of use and potential use, (4) the hydraulic character of the aquifer, (5) the areas of known ground-water contamination, and (6) the known locations of current and potential hydrocarbon, mineral, and geothermal resources in the Valley and Ridge and Blue Ridge provinces. Formation names used in this report are those of the Tennessee Division of Geology (Miller, 1974) and do not necessarily follow the usage of the U.S. Geological Survey.

GEOLOGY

The formations that make up the framework of the East Tennessee aquifer system range in age from Precambrian to Mississipian (table 1). They are composed of folded and faulted sedimentary rocks (limestones, shales, dolomites, sandstones, and conglomerate) in the Valley and Ridge physiographic province, and fractured sedimentary, metasedimentary, and crystalline igneous and metamorphic rocks of the Blue Ridge province. The rocks are overlain by a mantle of residual soil which in places may exceed 150 feet in thickness (De-Buchananne and Richardson, 1958). More commonly, however, the thickness of residual soil is less than 10 feet, and throughout the area it is not uncommon to see exposed rock with no soil. A veneer of alluvium, composed of boulders, gravel, silt, sand, and clay, covers the bottom of major valleys (Zurawski, 1979).

The structural setting of the East Tennessee aquifer system is very important because it is one of the major controlling influences on the occurrence of ground water, especially in the Valley and Ridge. The sedimentary rocks of the Valley and Ridge were folded and broken into a series of sheets that were thrust several miles northwestward. This deformation has resulted in a repetition of the same rock layers and a compartmentalization of aquifers (fig. 2). A map of the major structural features is shown in figure 3, and a section showing the generalized configurations of the rocks is shown in figure 4.

Toward the east in the Blue Ridge province, the rocks become progressivly more deformed and metamorphosed. Commonly, the rocks in this province are massive, and with the exception of the upper several hundred feet, are nonporous and impermeable. Within several hundred feet of land surface, fractures cut across the various rock types and provide homogeneous, secondary permeability.

The East Tennessee aquifer system is separated from other regional aquifers to the west by a zone of faulting. This zone occurs in a broad area which includes the eastern part of the Cumberland Plateau and the western part of the Valley and Ridge province. Faulting has generally occurred in the incompetent shales of the Rome Formation, causing repetition of the sequence of Rome Formation, Conasauga Group, and Knox Group through the Valley and Ridge province (fig. 2). These repeating sequences do not appear to be hydrologically continuous because of the impermeability of the faults and the basal shale which serves as the glide plane and accompanies the faulting.

The geology of East Tennessee has been studied in detail, and in addition to the more accessible references listed below, a store of detailed geologic information exists in quadrangle maps, geologic theses, and site reports that have not received widespread distribution but are nonetheless available. Of a more regional nature, the following publications were used for generalizing the geology presented in this report: Rodgers (1953); Neuman

(1955); Swingle (1959); King (1964); Neuman and Nelson (1965); LeGrand (1967); McMaster and Hubbard (1970); Harris and Milici (1977); Milici and Wedow (1977); and Milici, Hassis, and Statler (1979).

HYDROLOGY

The general hydrology of the Blue Ridge province is distinct from the Valley and Ridge province as shown by figures 5 and 6. Most of the water in the sedimentary, metasedimentary, and crystalline rocks of the Blue Ridge province occurs in the upper 200 feet, in interconnected fractures in the rock and in the pore spaces of overlying soil and regolith (fig. 5). Below several hundred feet, the weight of the overlying rock tends to keep the fractures closed, and regional ground-water flow below this depth is not considered to be significant.

Ground-water occurrence in the Blue Ridge is thus determined by the number, size, and degree of interconnection of the openings in the rocks and by the thickness of the saturated overburden (McMaster and Hubbard, 1970; and Zurawski, 1979). Ground-water circulation patterns tend to be localized rather than regional in extent, with relatively shallow flow paths (LeGrand, 1967). Recharge is areally distributed and discharge areas are local seeps, springs, and streams. Reported well yields and spring discharge are consistent with this interpretation, as is the water-quality distribution. It should be noted that few data exist from depths greater than 300 feet in the Blue Ridge.

In the Valley and Ridge province, it is known from records of water wells and other borings that solution cavities containing water are present at depths 900 to 1,000 feet below the surface (DeBuchanne and Richardson, 1956). Most solution openings, however, are confined to the upper 300 feet. Large spring discharges indicate a more active groundwater system at shallow depths than in the Blue Ridge. However, the highly variable well yields of the Valley and Ridge indicate this aquifer is more anisotropic and nonhomogeneous than the Blue Ridge province. In addition to solution cavities, ground water in the Valley and Ridge province occurs in fractures and, in some instances, along bedding planes of the carbonates and shales (fig. 6). The complexity of the structure and sparse data makes interpretation of the deep regional flow system not possible at this time.

In addition to the importance of structure in the Valley and Ridge province, rock type plays an important role in the hydrology. Carbonates are the most productive water-bearing formations in this area. According to DeBuchananne and Richardson (1956), many sinkholes and other karst features are common in the Valley and Ridge province where extensive solution of the underlying limestone and dolomite has taken place. In such areas, few surface streams are found; most of the drainage is through a well-developed underground drainage system, and the water table is likely to be deeper than in other areas.

There is evidence that solution is more extensive near perennial streams than elsewhere (DeBuchananne and Richardson, 1956). Industries close to rivers are more successful in obtaining large supplies of ground water than those in other locations. It is also likely that solution along zones of weakness in the rocks has determined the stream position in some areas.

Shales may be important water-bearing formations in the East Tennessee aquifer system, unlike in other areas of the State. Normally, shales have little effective primary porosity, and unless secondary openings are formed by fracturing, shales will yield little

water to wells. The rocks of East Tennessee have been folded and faulted extensively, however, and shales that are hard and brittle enough to support fractures are among the better aquifers of the area. Shales containing appreciable quantities of calcium carbonate yield more water than noncalcareous shales, as the fractures in such rock are susceptible to enlargement by the solvent action of water. In general, fractures in shale are more closely spaced than those in limestone and dolomite.

Sandstones and noncalcareous shales are composed of particles of minerals and rock more or less firmly cemented together. Rocks of these types found in East Tennessee contain practically no primary openings. Water is transmitted in secondary openings consisting of joints, fractures, and solution openings. Unlike limestone, dolomite, and calcareous shale, the openings in sandstone are not readily susceptible to dissolution by water. Sandstones and noncalcareous shales are not as widely distributed in East Tennessee as limestones, dolomites, and calcareous shales. However, rocks of this type, because of fracturing, will usually yield small supplies of water.

Recharge occurs by the percolation of rainfall through the residuum that overlies the East Tennessee aquifer system. Discharge occurs as springs, base flow to streams and rivers, and pumpage from wells. The residuum yields enough water to supply many domestic wells. During the late summer-early autumn, a period when water levels usually decline, many of these shallow wells may go dry. Water levels in this aquifer system fluctuate several feet in response to varying recharge and discharge conditions.

WATER QUALITY

The quality of water from the East Tennessee aquifer system is generally very good throughout its area of occurrence (fig. 7). The dissolved-solids concentrations in water from most wells were less than 250 mg/L. However, it should be noted that data are available from only one well with a depth greater than 500 feet.

Water from three wells on record had dissolved-solids concentrations of as much as 1,000 mg/L (table 2). Each of these occurrences was isolated and no discernible pattern was observed. None of the three wells was deeper than 135 feet below land surface; two were in a shale, and one was in a limestone. Such high concentrations of dissolved solids are local in extent, and may in part be caused by contamination. They do not reflect the regional water-quality trends, but they do point out that local anomalies are present.

The mode of occurrence of ground water in the Valley and Ridge province (fig. 6) makes contamination to this part of the aquifer a continuing problem. The highly anisotropic nature and occurrence of the water-bearing zones, the high permeability and rapid ground-water movement associated with the solution cavities in the folded carbonates, and the good quality and widespread utilization of the formations for drinking-water sources provide a combination of physical conditions that, on a regional scale, render the aquifer unsuitable for waste disposal. Water quality in the deeper formations of this aquifer system is not known, but dissolved-solids concentrations may be greater than 1,000 mg/L (fig. 6).

The quality of shallow ground water in the crystalline rocks of the Blue Ridge province is very good. Below the upper shallow flow system, however the rocks are effectively impermeable and nonporous.

In addition to much unpublished data, the following reports were used to compile information for this water-quality section: Glenn (1904); and DeBuchananne and Richardson (1956).

DRINKING-WATER SUPPLIES

The East Tennessee aquifer system is used extensively throughout its area of occurrence as an important source of drinking-water supplies (fig. 8). The yields are generally adequate for public and domestic supplies. Public water supplies from this aquifer system are listed in table 3. Little use has been made of water from depths greater than 500 feet. Below several hundred feet, ground water represents a resource whose quantity and quality are essentially unknown.

Most of the data for drinking-water supplies come from unpublished sources, primarily the Tennessee Department of Health and Environment. Historic use of water from this aquifer is documented in DeBuchananne and Richardson (1956); Swingle (1959); and Wilson and Johnson (1970).

CONTAMINATION

The East Tennessee aquifer system has 16 locations of documented contamination. The locations are shown in figure 9 and are described in table 4. Each occurrence of contamination is limited geographically and none is believed to pose an immediate threat to the aquifer except in localized areas.

HYDROCARBON, MINERAL AND GEOTHERMAL RESOURCE USE

The East Tennessee aquifer system includes many mineral deposits that were formed during several periods of Appalachian mountain building. These minerals are localized in two major mining areas, although numerous isolated deposits occur throughout East Tennessee. The occurrence of these deposits is generalized and shown in figure 10.

The Ducktown-Copperhill area of Polk County, in the extreme south-eastern part of the State, is the only copper mining area in the State. Copper sulfides occur in metamorphosed sediments of the Great Smoky Group. These deposits have been mined from the surface to a depth of about 2,500 feet.

The other major mining area is in the vicinity of Mascot and Jefferson City, in Knox and Jefferson Counties, where zinc and associated minerals are concentrated. In this area, zinc and lead sulfides occur in the carbonates of the Knox Formation. Other minerals that have been, or may possibly be mined, are gold, barite, galena, pyrite, and manganese.

Some potential for hydrocarbon resources exists throughout the Valley and Ridge province (fig. 10). The greatest potential probably exists along the western margin of the area, where the more deformed rocks of the Valley and Ridge province have buried a toe of Cumberland Plateau rocks. This buried toe is relatively undeformed and may contain hydrocarbons (Harris and Milici, 1977). Deep exploratory drilling for hydrocarbons is currently taking place in the Tennessee part of the Eastern Overthrust.

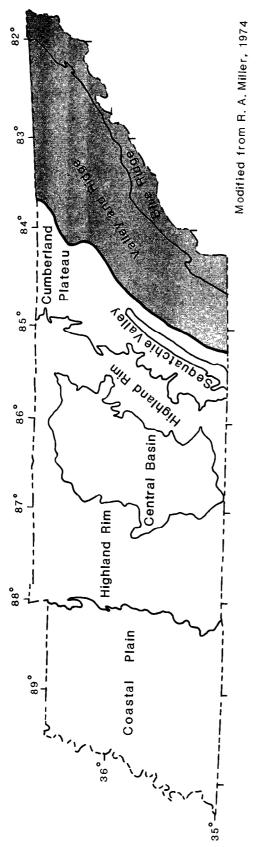
No geothermal resources are known to occur in the East Tennessee aquifer system.

SUMMARY

The East Tennessee aquifer system occurs in the Valley and Ridge and Blue Ridge physigraphic provinces. This aquifer system is composed of formations ranging in age from Precambrian to Mississippian. Limestone, dolomite, and calcareous shale are the principal water-bearing rocks of the area. Unlike the other regional aquifers, the East Tennessee aquifer system is delineated on the basis of its distinct structural and physiographic setting and not on its stratigraphy. Ground-water occurrence in this aquifer, particularly in the Valley and Ridge province, is unique because the water-bearing formations have been deformed by faulting and folding. Regional lateral flow in the permeable formations does not generally occur. For the most part, circulation is restricted to fractures that have been enlarged by solution. Faults that commonly occur within weak shale beds result in discontinuities that tend to isolate ground-water movement into discrete compartments. Ground-water conditions below a depth of about 300 feet are virtually unknown because of the structural complexity of the East Tennessee aquifer system and the paucity of data.

The East Tennessee aquifer system is classified as an underground drinking-water source under the criteria defined by the Safe Drinking Water Act. Water quality in the upper part of the aquifer is generally good to excellent, with dissolved-solids concentrations commonly less than 500 milligrams per liter. This aquifer system is used for drinking water throughout its area of occurrence in Tennessee. There are seven locations where contamination of the aquifer has been documented. However, these are limited geographically and none are thought to threaten the water quality of the aquifer on a regional basis.

Two main areas of mineral resource use occur within the East Tennessee aquifer system. Copper has been mined in the Ducktown-Copperhill area, and zinc and associated minerals are mined in Knox and Jefferson Counties. In addition to these two developed areas of mineral use, exploration for hydrocarbons is currently (1982) taking place along the eastern overthrust belt in the Valley and Ridge province.



0 25 50 75 MILES 0 25 50 75 KILOMETERS

EXPLANATION

Area of occurrence of the East Tennessee aquifer system

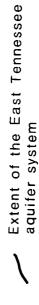


Figure 1.-- Areal extent of the East Tennessee aquifer system and physiographic provinces in Tennessee.

Table 1.--Hydrogeology of the formations comprising the East Tennessee aquifer system

| massive limestone member. Thickness 150 to 2,250 feet. In west, generally pure gray mascher. Parts containing some chart. Parts containing some chart. Parts contain some shaly beds. Shaly beds appear lower toward tree east and the formation becomes more snaly. Thickness 1,200 to 2,500 feet. Limestone, Siliceous, gray to bluish-gray, and shale with chert chart stringers. Thickness 100 to 250 feet. Shale, black fissile. The Chattanonga Shale is divided into three members. The thickness long freet. Thick beds of limestone and dolomite. The majority of these beds mite. The majority of these beds mite. The majority of these beds mate. The majority of these beds are sandy but a few are cherty. Thickness is generally less than 300 feet. Largely greenish to brownish shale and oeds of siltstone and limeastone. Hematite beds encountered at varying depths. Thickness |
|--|
| Thickbedded to massive, well- cemented quartz sandstone. Nedium- to coarse-texture. Nudstone and limestone with some sand, shale, and silty limestone. The limestone is more calcareous. Thickness 200 to 400 feet. Bluish-gray, well bedded or platy to nodular limestone with inter- bedded shaly partings. Few thin beds of volcanic ash present. Many fossils in formation. Thickness approximately 2,000 feet. Dolomite, gray and brown, fine- grained to granular, and dense grained to granular, and dense |
| |

| | <u> </u> | | | T | T | , MC | |
|---|---|--|---|---|---|---|--|
| Several gallons per minute yielgs to domestic wells. Springs flow as much as 450 gallons per minute. | Small to moderately large yields. | Small to moderately large yields. | Small to moderately large yields. | Small to moderately large yields. | Small to moderately large yields. | Yields are usually low, generally less than several gallons per minute. | Small to moderate Yields. Only one of Six inventoried springs had an esti- mated yield greater than 100 gallons per minute. |
| Ground water occurs in fractures in shale and sandstone and in solution channels in the dolomite. The upper zone is more permeable than the lower part of the formation. | Ground water limited to fractures, joints, and bedding planes. Highly variable porosity and permeability. Rock has massive nonporous matrix. | Most ground water occurs only Small to moderately in zones of secondary porosity large yields. and permeability. | Ground water restricted to fractures in the upper 200 feet of land surface. Most ground water occurs in zones of secondary porosity and permeability. | Ground water restricted to fractures in the upper 200 feet of land surface. | Ground water occurs in zones of secondary porosity and permeability in the upper 200 feet of land surface. | Ground water restricted to fractures which occur in the upper 200 feet of land surface. | Ground water restricted to fractures in the upper 200 feet of land surface. |
| Widespread occurrence in Valley and Ridge. North- east to southwest linear outcrops. Repeated occur- rence due to faulting. | widespread occurrence in Valley and Ridge. North- east to southwest linear outcrops. Repeated occur- rence due to faulting. | Widespread occurrence in Valley and Ridge. North- east to southwest Inear outcrops. Repeated occur- rence due to faulting. | Widespread occurrence in Valley and Ridge. North- east to southwest linear outcrops. Repeated occur- rence due to faulting. | Widespread occurrence in Valley and Ridge. Wortheast to Southwest linear outcrops. Repeated occurrence due to faulting. | Widespread occurrence in Valley and Ridge. North east to southwest linear outcrops. Repeated occur- rence due tofaulting. | Widespread occurrence in Valley and Ridge. Northeast to southwest linear outcrops. Repeated occurrence due to faulting. | Widespread occurrence in Valley and Ridge. North- east to southwest linear outcrops. Repeated occur- rence due to faulting. |
| Sandstone, siltstone, shale, dolo- mite, and limestone. Shale and siltstone predominate with promi- nant sandstone beds. In south- east dolomite constitutes half the formation. Thickness varies from 200 to 1,500 feet. | Dolomite, blue-gray to light-gray, silty. Limestone present in lower part and sandy beds occur near the base. Thin layer of argillaceous, shaly dolomite in upper part. Chert present throughout. Thickness approximately 1,000 feet. | Sandstone and quartzite, fine- grainea. Gray to greenish, witn shale. Barely exceeds 100 feet in thickness. | Sandstone, white, quartzite cemented. Medium- to coarse- grained. Commonly occurs in ledges. Sandstone is inter- bedded with dark green silty, sandy, or clay shale mixed with yery fine siltstones and sand- stones. Thickness about 600 feet. | Shale, silty, sandy, dull green to brown, micaceous. Thickness approximately 500 feet. | Quartzite, medium-bedded, fine- grain, white, vitreous, in part feldspathic. Approximately 250 feet thick. | Shale, silty, sandy, containing flakes of detrital mica. Lenses of sandstone present but are relatively thin. Thickness 800 feet. | Conglowerate, gray, peobly arkose, siltstone, and shale. Irregular bedding, micaceous arkose and shale near midgle and base. Thickness about 1,200 feet. |
| Rome Formation | Shady Dolomite | Helenmode Formation | Hesse Sandstone | Murray Shale | Nebo Sandstone | Nichols Shale | Cochran Formation |
| <u></u> | | | CAMBRIAN | | | | |

Table i.--Hydrogeology of the formations comprising the East Tennessee aquifer system--Continued

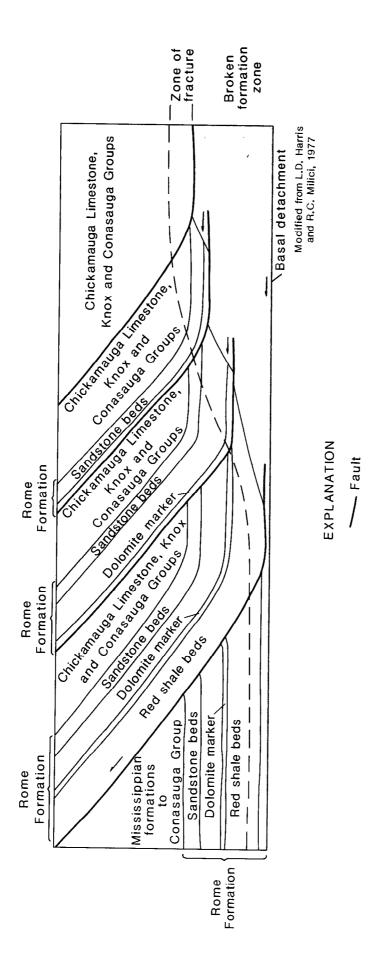


Figure 2.-- Generalized cross section of a fault block in the Valley and Ridge province showing repetition of formations.

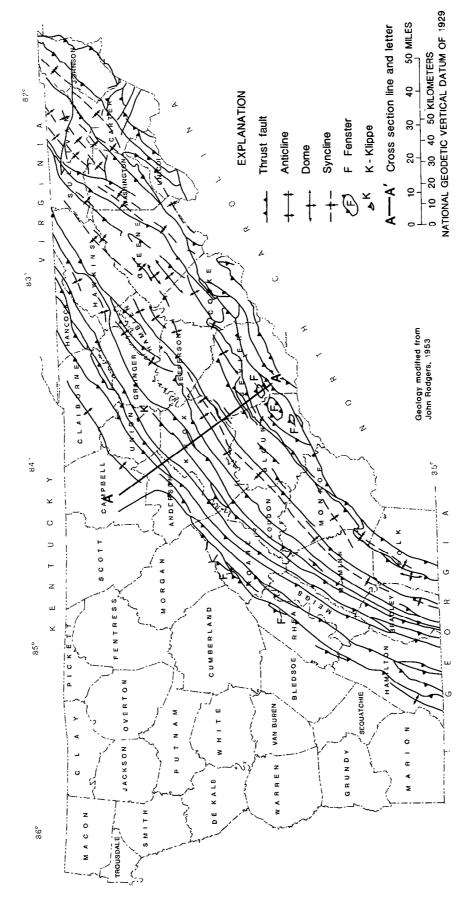


Figure 3.-- Structural features of East Tennessee aquifer system.



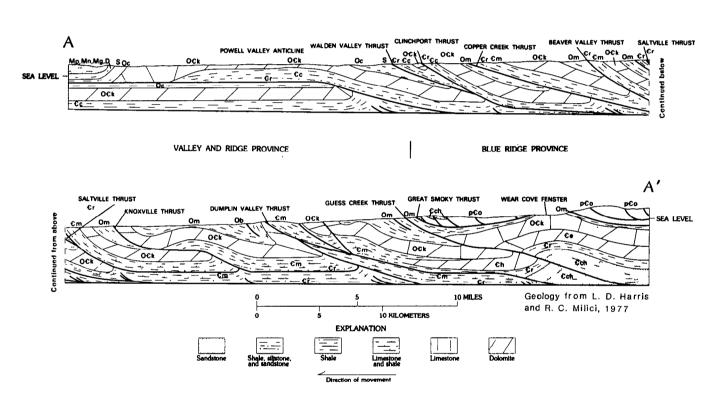


Figure 4.--Generalized geologic cross section of East Tennessee.

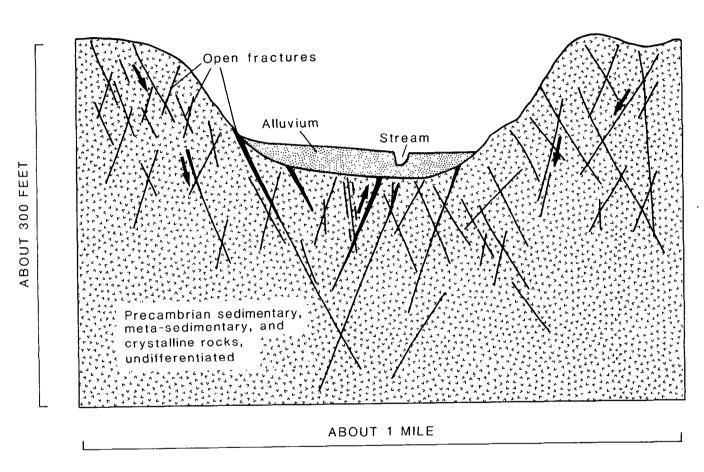


Figure 5.-- Conceptual model of ground-water occurrence in the Blue Ridge province.

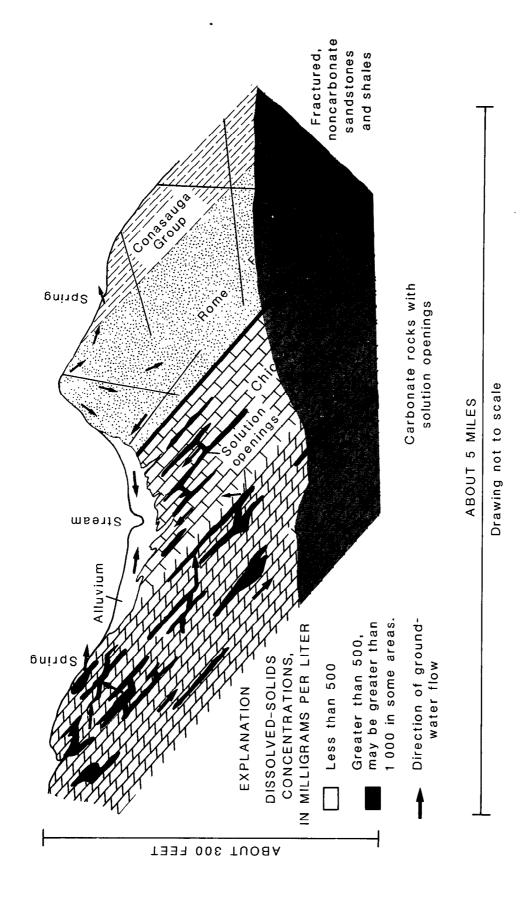


Figure 6.-- Conceptual model of ground-water occurrence and generalized water quality in the Valley and Ridge province.

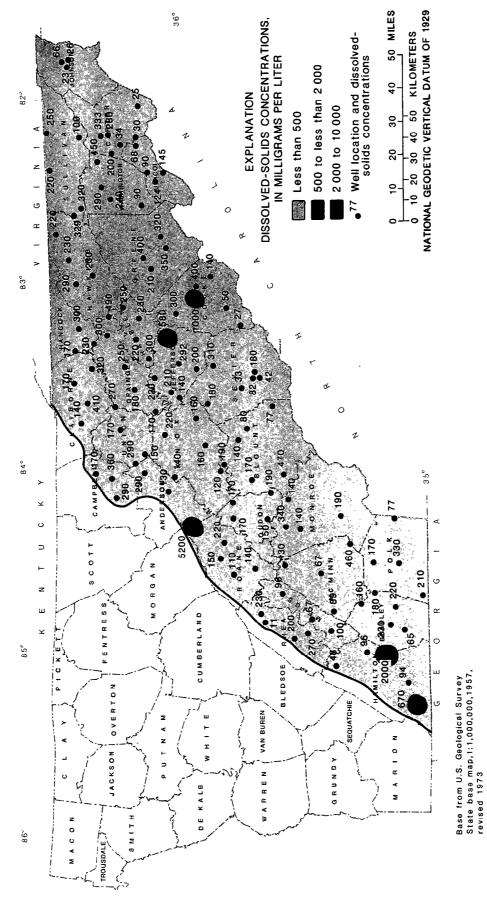


Figure 7.-- Dissolved-solids concentrations in the East Tennessee aquifer system.

Table 2.--Dissolved-solids concentrations in water from the East Tennessee aquifer system

[Data source codes: 1, DeBuchananne and Richardson (1956); 2, McMaster and Hubbard (1970); 3, Maclay (1962); 4, Hollyday and Goddard (1979); 5, Zurawski (1979); 6, Unpublished U.S. Geological Survey records; E Estimated from specific conductance]

| County | Location | Well depth, in feet | Water-bearing of formation | Dissolved solids, concentra- concentra- cions, in milligrams per liter | Data source |
|-----------|---|---|--|--|----------------------------|
| Anderson | Andersonville 0.5 mi NE Clinton 0.5 mi W | 114 Spring | Chickamauga Limeston Knox Group | ne 290 E 130 E | 1 |
| Blount | Friendsville 3 mi SE Mentor 3 mi N Rockford 0.5 mi S Tallassee 4.5 mi N Walland 2.5 mi N Tremont | Spring 264 460 64 77 130 | Knox Group Holston Formation Lenoir Limestone Athens Shale do | 170 E 190 E 140 E 470 E 80 E 77 | 1 1 1 1 1 2 |
| Bradley | Benton 4.5 mi NW Charleston 2.5 mi SW Cleveland 1 mi SW McDonald Ocoee 4 mi W | 100 Spring 423 30 95 | Conasauga Group do do do do | 180 E 160 E 230 E 65 E 220 E | 1 1 1 1 1 |
| Campbell | Duff 3 mi SE Jacksboro 1 mi E Lafollette 4.5 mi SE | 230 4219 300 | Chickamauga Limestor Newala Formation Copper Ridge Dolomite. | 170 E 290 E 360 E |]]] |
| Carter | Elizabethton 3 mi S Elizabethton Hampton 2 mi SW Milligan College 0.5 mi S Shell Creek 1 mi SW Unicoi 6.5 mi E | 135 95 109 Spring Spring | Honaker Dolomite do Shady Dolomite Knox Group Precambrian crys- talline complex. do | 280 333 34 E 210 E 25 E | 3 3 1 1 1 |
| Claiborne | Clouds 3.5 mi S Goin 4 mi NW Tazewell 3 mi NE Thorn Hill 4.5 mi NW | Spring Spring Spring 128 | Longview Dolomite Copper Ridge Dolomite. Mascot Dolomite Conasauga Group | 410 E 140 E 170 E 320 E | 1 1 1 |

Table 2.--Dissolved-solids concentrations in water from the East Tennessee aquifer system--Continued

| County | Location | Well depth, in feet | Water-bearing of formation | Dissolved solids, concentra- tions, in milligrams per liter | Data source |
|----------|---|-------------------------------------|--|---|-----------------------|
| Cocke | Bybee French Broad 1 mi S Hartford 3.5 mi NW Newport 1.5 mi NE Parrottsville 4.5 mi SE Indian Camp Creek | 47 51 15 135 105 194 | Sevier Shale Sandsuck Shale Shady Dolomite Sevier Shale Honaker Dolomite | 300 E 140 E 50 E 1000 E 400 E 27 |]]]] 2 |
| Grainger | Blaine Joppa 2 mi E Mooresburg 3 mi W | 75 Spring Spring | Conasauga Group Copper Ridge Dolomite. Chickamauga Limesto | | 1 |
| | Rutledge 3.5 mi E | 210 | Copper Ridge Dolomite. | 250 E | Ī |
| Greene | Cedar Creek 6 mi NE | 25 | Knox Group | 350 E |] |
| | Greenville 2.5 mi NW Mosheim 3 mi SE | 310 90 | Sevier Shale do | 400 E 210 E |]] |
| | Tusculum College 5 mi S | 195 | Knox Group | 320 E | 1 |
| Hamblen | Morristown 1.5 mi NW | Spring | Newala Formation | 220 E | 1 |
| | Russellville 3.5 mi S | Spring | | 230 E |] |
| | Talbott 1 mi N Whitesburg | 201 Spring | Newala Formation Knox Group | 300 E 250 E |] |
| Hamilton | Chattanooga | 65 | Knox Group | 670 E | 7 |
| | Georgetown 5 mi SW McDonald 5 mi NW | Spring 67 | do Chickamauga Limestone. | 96 E 2000 E |] |
| | Sale Creek 0.5 mi E Tyner 2 mi W | 60 200 | Newman Limestone Newala Formation | 48 E 94 E | 1 |
| Hancock | Luther 6.5 mi W | 45 | Pumpkin Valley Shal | | 1 |
| | Thorn Hill 9.5 mi N | Spring | Newman Limestone | 170 E | 1 |
| | Thorn Hill 5 mi NE | 43 | Chickamauga Limesto | ne 230 E | 1 |
| Hawkins | Church Hill 2.5 mi N | Spring | Conasauga Group | 220 E | 7 |
| | Eidson 3 mi SE | 17 | Newman Limestone | 290 E | 1 |
| | Mooresburg 2.5 mi_E | 220 | Moccasin Formation | 490 E |] |
| | Rogersville 3 mi E | Spring | Copper Ridge Dolomite. | 230 E | 1 |
| | Surgoinsville 2 mi N | Spring | Conasauga Group | 230 E | 7 |

Table 2.--Dissolved-solids concentrations in water from the East Tennessee aquifer system--Continued

| County | Location | Well depth, in feet | Water-bearing of formation t | dissolved solids, concentrations, in alligrams per liter | source |
|-----------|--|----------------------------------|--|--|------------------|
| Jefferson | Dandridge 0.25 mi NW Dandridge 5.5 mi SW Jefferson City 3 mi NW | 400 117 Spring | Copper Ridge/Che- pultepec Dolomite. Sevier Shale Copper Ridge | 292 200 E 220 E | 4 1 1 |
| | New Market 6.5 mi SW Strawberry Plains 0.5 mi E White Pine 1.5 mi SW | 130 Spring 105 | Dolomite. Mascot Dolomite Lenoir Limestone Knox Group | 140 E 210 E 580 E | 1 1 1 |
| Johnson | Mountain City 1 mi W Mountain City 1 mi E Mountain City 1.5 mi NE | 107 Spring Spring | Rome Formation Shady Dolomite Rome Formation | 23 E 26 E 66 E |]]] |
| Knox | Corryton 4 mi SW Heiskell 0.5 mi W Knoxville 2.5 mi SE Louisville 4 mi N Mascot 5.5 mi S | Spring 60 168 30 168 | Chickamauga Limeston do Holston Formation Chepultepec Dolomite Mascot Dolomite | 140 E 160 E | 1 1 1 1 |
| Loudon | Greenback Lenoir City 2.5 mi NW Loudon 5 mi SE Martel | 82 68 Spring | Knox Group Chickamauga Limeston Copper Ridge Dolomit Lenoir Limestone | | 1 1 1 |
| McMinn | Athens Big Spring 5.5 mi E Erie 3 mi SE Etowah 4 mi E | Spring Spring 26 | Kingsport Formation Knox Group Longview Dolomite Athens Shale | 67 E 89 E 130 E 460 E | 1 1 1 |
| Meigs | Big Spring Decatur 3.5 mi SW Ten Mile 4.5 mi SW | Spring Spring 54 | Chickamauga Limeston Knox Group do | e 100 E 67 E 96 E |]]] |
| Monroe | Madisonville 0.5 mi E Philadelphia 5.5 mi SE Tellico Plains Vonore 2.5 mi E | 80 85 90 300 | Conasauga Group Newala Formation Shady Dolomite Knox Group | 140 E 340 E 190 E 140 E |]]] |
| Polk | Archville 1.5 mi SW Conasauga Delano 2.5 mi S Turtletown 1.5 mi E | 200 125 Spring 60 | Ocoee series Athens Shale Conasauga Group Great Smokey conglomerate. | 330 E 210 E 170 E 77 E |]]] |

Table 2.--Dissolved-solids concentrations in water from the East Tennessee aquifer system--Continued

| County | Location | Well depth, in feet | Water-bearing of formation | Dissolved solids, concentra-tions, in milligrams per liter | Data source |
|------------|--|---|--|--|-----------------------|
| Rhea | Evensville Evensville 4 mi SE Grandview 2.5 mi SE Spring City 1 mi S | 84 85 43 25 | Chickamauga Limesto do Knox Group do | ne 200 E 270 E 230 E 11 E |]]] |
| Roane | Erie 6 mi N Kingston 2.5 mi NW Kingston 6 mi SW Kingston 4 mi E Oak Ridge | 45 12 69 88 90 | Chickamauga Limesto Conasauga Group Chickamauga Limesto Conasauga Group do | 150 E | 1 1 1 1 6 |
| Sevier | Boyds Creek 1 mi W Gatlinburg Gatlinburg Gatlinburg Pigeon Forge 2.5 mi SW Sevierville 6.5 mi NE | Spring 100 255 230 36 38 | Knox Group Great Smokey conglomerate. Snowbird Group do Sandsuck Shale Sevier Shale | 180 E 180 E 82 42 33 E 310 E |] 5 5 1 1 |
| Sullivan | Blountville 4 mi NW Bluff City 4.5 mi SE Bristol Fall Branch 3.5 mi N | 209 Spring 280 80 | Knox Group Sevier Shale Knox Group Sevier Shale | 220 E 100 E 250 E 330 E |]]] |
| Unicoi | Erwin Erwin 3 mi SW Erwin 4 mi S Unicoi 5 mi E | 135 122 Spring 30 | Honaker Dolomite Erwin Formation Unicoi Formation Shady Dolomite | 90 124 145 68 E |]]] |
| Union | Andersonville 6 mi E Andersonville 4 mi NE Maynardville 3.5 mi N Powder Springs 4 mi N | Spring 350 Spring 20 | Kingsport Formation Chickamauga Limesto Ottosee Shale Conasauga Group | |]]]] |
| Washington | Johnson City 6 mi NW Jonesboro 3.5 mi W Washington College 4 mi Watauga 3 mi W | 342 Spring S 57 136 | Knox Group do do Sevier Shale | 290 E 240 E 90 E 450 E |]]] |

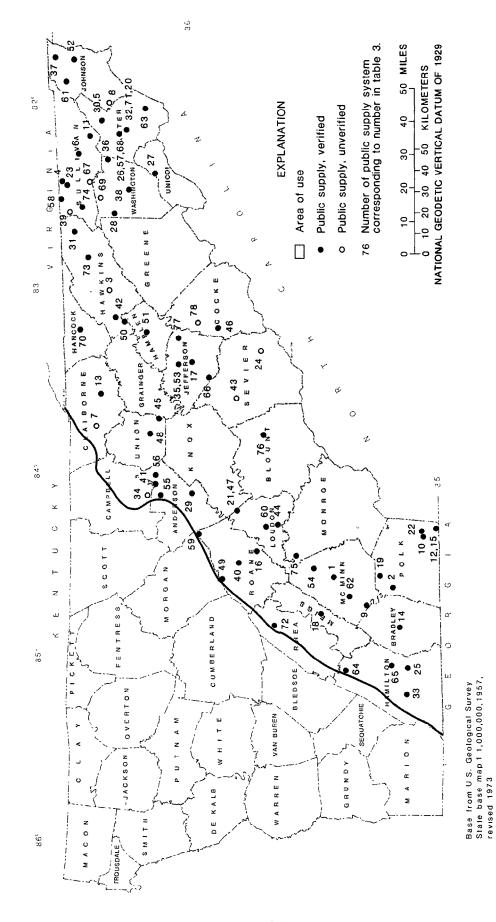


Figure 8.-- Public-supply systems and area of use of water from the East Tennessee aquifer system.

Table 3.--Summary of public-supply systems using water from the East Tennessee aquifer system

[Data source codes: 1, Reported - Tennessee Division of Water Resources; 2, Reported - Tennessee Division of Water Quality Control; 3, Tennessee comprehensive joint water and related land resources planning, Tennessee Division of Water Resources]

| Location No. | System | County | Data source |
|------------------|------------------------------------|------------|----------------|
| 1 | Athens | McMinn | 1,2,3 |
| 2 | Benton | Polk | 1,2,3 |
| 2 3 | Big Creek U.D. | Hawkins | 2 |
| 4 5 6 7 | Bloomingdale U.D. | Sullivan | 1,3 |
| 5 | Blue Springs U.D. | Carter | 2 |
| 6 | Bluff City | Sullivan | 1,2,3 |
| 7 | Cape Norris Subdivision | Claiborne | 2 |
| 8 9 | Carderview U.D. | Johnson | 2 |
| | Charleston-Calhoun U.D. | McMinn | 1,2,3 |
| 10 | Cherokee Hills | Polk | 1,2,3 |
| 11 | Chinquapin Grove U.D. | Sullivan | 1,2,3 |
| 12 | Cities Service | Polk | 1,2,3 |
| 13 | Claiborne Co. U.D. | Claiborne | 1,2,3 |
| 14 | Cleveland | Bradley | 2,3 |
| 15 | Copperhill | Polk | 1,2,3 |
| 16 | Cumberland U.D. | Roane | 1,2,3 |
| 17 | Dandridge | Jefferson | 1,2,3 |
| 18 | Decatur | Meigs | 1,2,3 |
| 19 | Delano | Po1k | 1,2,3 |
| 20 | Dividing Ridge Utilities, Inc. | Carter | 2 |
| 21 | Dixie Lee U.D. | Loudon | 1,2,3 |
| 22 | Ducktown | Polk | 2,3 |
| 23 | East Kingsport U.D. | Sullivan | 1,3 |
| 24 | East Sevier U.D. | Sevier | 2 |
| 25 | Eastside U.D. | Hamilton | 1,2,3 |
| 26 | Elizabethton | Carter | 1,2,3 |
| 27 | Erwin | Unicoi | 1,2,3 |
| 28 | Fall Branch | Washington | 1,2,3 |
| 29 | First U.D. of Anderson Co. | Anderson | 1,2,3 |
| 30 | First U.D. of Carter Co. | Carter | 1,2,3 |
| 31 | First U.D. of Hawkins Co. | Hawkins | 1,2,3 |
| 32 | Hampton U.D. | Carter | 1,2,3 |
| 33 | Hixson U.D. | Hamilton | 1,2,3 |
| 34 | Indian River | Campbell | 2 |
| 35 | Jefferson City | Jefferson | 1,2,3 |
| 36 | Johnson City | Washington | 1,2,3 |
| 37 | Johnson Co. Utilities Nos. 1 and 2 | Johnson | 2,3 |
| 38 | Jonesboro | Washington | 1,2,3 |
| 39 | Kingsport | Sullivan | 2 |

Table 3.--Summary of public-supply systems using water from the East Tennessee aquifer system--Continued

| Location No. | System | County | Data source |
|-----------------|--------------------------------|------------|----------------|
| 40 | Kingston | Roane | 1,2,3 |
| 41 | Lake City | Anderson | 1,2,3 |
| 42 | Lakemont | Hawkins | 1,2,3 |
| 43 | Little Ponderosa | Sevier | 2 |
| 44 | Loudon | Loudon | 1,2,3 |
| 45 | Luttrell-Blaine-Corryton U.D. | Union | 1,2,3 |
| 46 | L.W. Hooper | Cocke | 1,2,3 |
| 47 | Martel U.D. | Loudon | 2 |
| 48 | Maynardville | Union | 1,2,3 |
| 49 | Midtown Water Co. | Roane | 1,3 |
| 50 | Mooresburg U.D. | Hawkins | 1 , 2,3 |
| 51 | Morristown | Hamblen | 1,2,3 |
| 52 | Mountain City | Johnson | 1,2,3 |
| 53 | New Market U.D. | Jefferson | 2 |
| 54 | Niota | McMinn | 1,2,3 |
| 55 | Norris | Anderson | 1,2,3 |
| 56 | North Anderson Co. U.D. | Anderson | 1,2,3 |
| 57 | North Elizabethton Water Co-op | Carter | 2 |
| 58 | North Kingsport U.D. | Sullivan | 1,3 |
| 59 | Oliver Springs | Roane | 1,2,3 |
| 60 | Piney | Loudon | 1,2,3 |
| 61 | Pleasant Valley U.D. | Johnson | 1,2,3 |
| 62 | Riceville U.D. | McMinn | 1,2,3 |
| 63 | Roan Mountain Water Co. | Carter | 1,2,3 |
| 64 | Sale Creek | Hamilton | 1.2.3 |
| 65 | Savannah Valley U.D. | Hamilton | 2 |
| 66 | Shady Grove U.D. | Jefferson | 2 |
| 67 | Sharps Creek Subdivision | Sullivan | 2 |
| 68 | Siam U.D. | Carter | $\bar{2}$ |
| 69 | Sinking Creek Spring | Washington | 2 2 2 2 1,2,3 |
| 70 | Sneedville U.D. | Hancock | 1.2.3 |
| 71 71 | South Elizabethton U.D. | Carter | 2 |
| 72 | Spring City | Rhea | 1,2,3 |
| 73 | Surgoinsville U.D. | Hawkins | 1,2,3 |
| 74 74 | Sullivan Gardens U.D. | Sullivan | 1,2,3 |
| 75 | Sweetwater | Monroe | 2 |
| 76 76 | Walland | Blound | 1,2,3 |
| 70 77 | White Pine | Jefferson | 1,2,3 |
| 77 78 | Wood Acres Subdivision | Cocke | 2 |

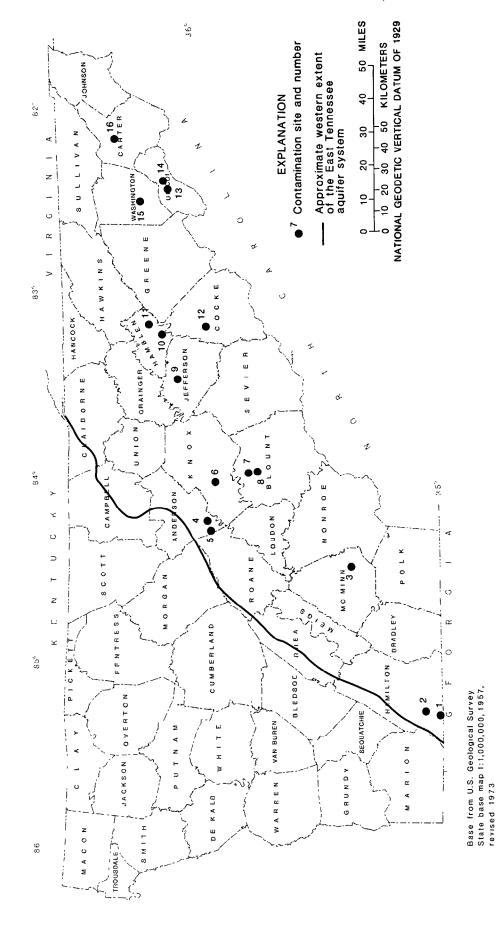


Figure 9.-- Contamination sites in the East Tennessee aquifer system.

Table 4.--Description of contamination sites

[Documentation Codes: a, Tennessee Division of Water Quality Control, unpublished records; b, Webster (1976); c, Residual Waste Study, Tennessee Division of Water Quality Control; d, Hyfantis (1980)]

| Site identification No. | Location | Type of contamination | Documentation | Stratigraphic interval | Comments |
|-------------------------------|-----------------------------------|------------------------------------|---------------|---------------------------|--|
| - | Hamilton County, Residue Hill. | Industrial wastes | го | Knox Group | Initial water-quality data were collected from eight monitoring wells around Residue Hill. Data indicated ground-water contamination by phenolic compounds, chlorinated hydrocarbons, several metals, and several volatile organics such as benzene and toluene. |
| 2 | Hamilton County, Chattanooga. | Industrial wastes | ď | Knox Group | A test well, drilled by TVA, encountered an oily, organic substance with an odor like coal tarcreosote. A coal gasification plant is said to have been located nearby, and the area served as an industrial dump for years. |
| ო | McMinn County | Industrial wastes | rs | | Ground water and surface water were contaminated with iron, manganese, COD, oil, grease, phenols, and diphenyl ether by dumping finishing oils into trenches. Wastewater from lagoons drained through solution openings in the limestone to a nearby creek. |
| 4 | Anderson County | Low level radio- active wastes. | Ф | Conasauga Group. | Ongoing program to define exact areas and constituents. No present evidence of extensive migration of contaminants. |
| ഹ | Anderson County | Industrial wastes | ત્ય | Conasauga Group. | Nitrate, mercury, and other industrial wastes have contaminated area ground and surface water. |

Table 4.--Description of contamination sites--Continued

| Site | | | | | |
|-----------------------|-------------------------------|-------------------------------------|---------------|---------------------------|--|
| identification No. | Location | Type of contamination | Documentation | Stratigraphic interval | Comments |
| 9 | Knox County | Industrial wastes | ro | Knox Group | Indiscriminate disposal of indus- trial wastes resulted in manga- nese contamination of ground water. |
| 7 | Blount County | Industrial wastes | rg . | Conasauga ? Group. | Earthen pits were used for the disposal of oily wastes. A nearby spring was found to be contaminated with oil and grease. |
| ω | Blount County | Industrial wastes | rt . | Conasauga ? Group. | This facility has been used for the disposal of fluoride dust from air pollution control facil- ities. Ground water in the area was contaminated by fluoride. |
| 6 | Jefferson County | | ro T | Knox Group | Mining operations have resulted in significant increases of zinc and, at times, suspended solids and turbidity levels in ground water. |
| 10 | Hamblen County | Industrial wastes | U | Knox ? Group | Objectionable quantities of organic compounds, dissolved solids, iron, manganese, sodium, sulfate, and phenols. Dissolved solids, sulfates, and phenols exceed drinking-water standards. Extent of ground-water degradation in the vicinity undetermined, but there is a high potential for continued, widespread degradation. |
| F | Morristown- Hamblen County | Municipal and industrial wastes? | U | Knox ? Group | High iron and manganese concentrations exceed drinking-water standards, definitely associated with the landfill. Hardness and dissolved solids higher than normal for the area, but cannot be specifically linked to the landfill at this time. |

Table 4.--Description of contamination sites--Continued

| Site identification No. | Location | Type of contamination | Documentation | Stratigraphic interval | Comments |
|-------------------------------|---------------------------------|---------------------------------------|---------------|---------------------------|--|
| 12 | Cocke County, Newport. | Laboratory wastewater | res | Knox Group | Laboratory wastewater was discharged into a sinkhole which resulted in the degradation of the ground-water quality. Studies indicated severe degradation of the ground water in the area of a spray irrigation system. |
| 13 | Unicoi County, Bumpass Cove. | Hazardous wastes | 9 | Shady Dolomite. | The illegal dumping of hazardous waste into an approved sanitary landfill resulted in the contamination of area ground water. Methylene chloride and trichloroethylene were found in a resident's well. |
| 14 | Unicoi County, Erwin. | Industrial and radiological wastes | Ф | Honaker ? Dolomite. | The disposal of industrial and radiological wastes has resulted in the contamination of ground water locally. |
| 15 | Washington County, Telford. | Industrial wastes | יט | Knox Group | The discharge of industrial wastes into an unlined earthen pond resulted in the contamination of surface and ground water by fluoride and nitrate. |
| 91 | Carter County, Elizabethton. | Industrial wastes | ro | Homaker Dolomite. | Waste disposal and solid residues have caused ground-water contam- ination by copper and zinc. |
| | | | | | |

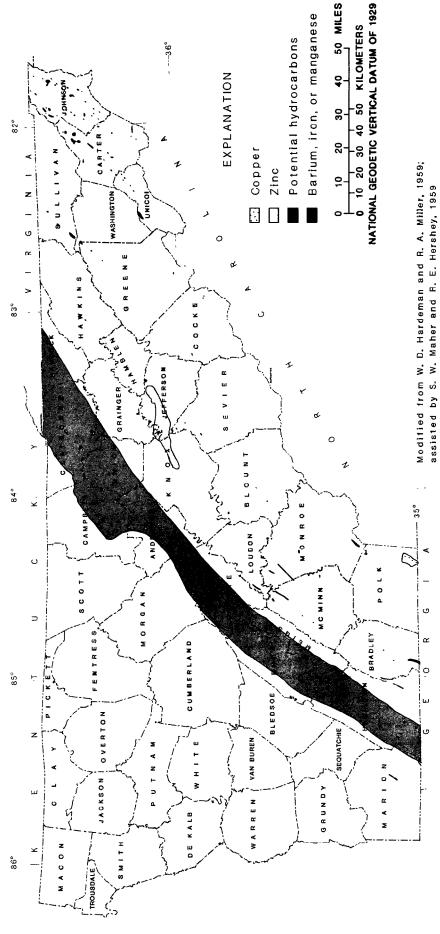


Figure 10.--Current and potential hydrocarbon, mineral, and geothermal resources.

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